# Assimilating Irregularly Spaced Sparsely Observed Turbulent Signals with Hierarchical Bayesian Reduced Stochastic Filters

#### Kristen Brown

Department of Mathematics North Carolina State University

Collaborator: John Harlim (NC State)

November 15, 2012

• Observations from nature are often noisy, temporally irregular and spatially sparse. On the other hand, the typical predictor model is resolved on regularly spaced grid points.

- Observations from nature are often noisy, temporally irregular and spatially sparse. On the other hand, the typical predictor model is resolved on regularly spaced grid points.
- $\bullet$  Typically, various interpolation techniques are used to merge multiple satellite data

For example: cloud clearing algorithm [Chahine, 1973].

- Observations from nature are often noisy, temporally irregular and spatially sparse. On the other hand, the typical predictor model is resolved on regularly spaced grid points.
- Typically, various interpolation techniques are used to merge multiple satellite data

For example: cloud clearing algorithm [Chahine, 1973].

• Processing data is often unavoidable For example: estimating rainfall given rain-guage measurements.

- Observations from nature are often noisy, temporally irregular and spatially sparse. On the other hand, the typical predictor model is resolved on regularly spaced grid points.
- Typically, various interpolation techniques are used to merge multiple satellite data

For example: cloud clearing algorithm [Chahine, 1973].

- Processing data is often unavoidable For example: estimating rainfall given rain-guage measurements.
- The goal of this talk is to asses the effect of processed data assimilated in the presence of model error.

- Observations from nature are often noisy, temporally irregular and spatially sparse. On the other hand, the typical predictor model is resolved on regularly spaced grid points.
- Typically, various interpolation techniques are used to merge multiple satellite data

For example: cloud clearing algorithm [Chahine, 1973].

- Processing data is often unavoidable For example: estimating rainfall given rain-guage measurements.
- The goal of this talk is to asses the effect of processed data assimilated in the presence of model error.
- In particular, we will use hierarchical Bayesian framework.

#### Standard Bayesian Approach

#### Definition

- U : Random variable of the model state
- $ightharpoonup ilde{V}$  : Random variable of the irregularly spaced observations
- $ightharpoonup P(U, ilde{V})$  : Joint density between two random variables

# Standard Bayesian Approach

#### Definition

- U : Random variable of the model state
- $ightharpoonup ilde{V}$  : Random variable of the irregularly spaced observations
- $ightharpoonup P(U, \tilde{V})$  : Joint density between two random variables

Canonical Discrete-Time Filtering Problem:

$$u_{m+1} = f(u_m) + \sigma_{m+1}, \quad \sigma \sim \mathcal{N}(0, r),$$
  
 $\tilde{v}_m = g(u_m) + \epsilon_m, \quad \epsilon_m \sim \mathcal{N}(0, r^\circ)$ 

# Standard Bayesian Approach

#### Definition

- U: Random variable of the model state
- $ightharpoonup ilde{V}$  : Random variable of the irregularly spaced observations
- $ightharpoonup P(U, \tilde{V})$  : Joint density between two random variables

Canonical Discrete-Time Filtering Problem:

$$u_{m+1} = f(u_m) + \sigma_{m+1}, \quad \sigma \sim \mathcal{N}(0, r),$$
 
$$\tilde{v}_m = g(u_m) + \epsilon_m, \quad \epsilon_m \sim \mathcal{N}(0, r^\circ)$$

Solution: Apply the Bayesian Theorem:

$$P(U|\tilde{V}) \propto P(U)P(\tilde{V}|U).$$

#### Hierarchical Bayesian Approach

Consider  $v \in V$  to be the random variable of interpolated observations at the regular model grid points. Our approach is to apply

$$P(U|\tilde{V}, V) \propto P(U)P(V|\tilde{V}, U)$$
  
  $\propto P(U)P(\tilde{V}|U, V)P(V|U).$ 

# Hierarchical Bayesian Approach

Consider  $v \in V$  to be the random variable of interpolated observations at the regular model grid points. Our approach is to apply

$$P(U|\tilde{V},V) \propto P(U)P(V|\tilde{V},U)$$

$$\propto P(U)P(\tilde{V}|U,V)P(V|U).$$

Step 1: We apply  $P(\tilde{V}|U,V)P(V|U)$  through an interpolation to obtain  $P(U)P(V|\tilde{V},U)$ . We compare a statistical interpolation called kriging with a deterministic linear interpolation.

# Hierarchical Bayesian Approach

Consider  $v \in V$  to be the random variable of interpolated observations at the regular model grid points. Our approach is to apply

$$P(U|\tilde{V},V) \propto P(U)P(V|\tilde{V},U)$$

$$\propto P(U)P(\tilde{V}|U,V)P(V|U).$$

Step 1: We apply  $P(\tilde{V}|U,V)P(V|U)$  through an interpolation to obtain  $P(U)P(V|\tilde{V},U)$ . We compare a statistical interpolation called kriging with a deterministic linear interpolation.

Step 2: We apply  $P(U)P(V|\tilde{V},U)$  through a reduced stochastic Fourier based filter.

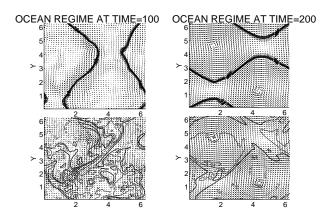
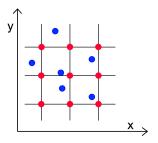


Figure: The 2 layer QG model with baroclinic instability, resolved with  $128 \times 128$  grid points in a 2D periodic domain [Smith et al, 2002]. The radius of deformation is chosen to mimic ocean turbulence. The top panels show the barotropic velocity field (arrows) and streamfunction  $\Psi$ , (contour) and the bottom panels show the baroclinic velocity field and streamfunction  $\Psi$  (bottom) at two different times.

Given two-dimensional noisy, sparse observations from the solution to the two-layer quasi-geostrophic model with baroclinic instability, the first task is to interpolate to a regular  $6\times 6$  grid.



Kriging is a maximum likelihood estimator of a random field Z modeled by

$$Z(s) = \mu(s) + \delta(s),$$

assuming Gaussian, stationary noises  $\delta(s) \sim \mathcal{N}(0, C(s, s))$ .

Kriging is a maximum likelihood estimator of a random field Z modeled by

$$Z(s) = \mu(s) + \delta(s),$$

assuming Gaussian, stationary noises  $\delta(s) \sim \mathcal{N}(0, C(s, s))$ .

The steps of kriging:

- 1. Estimate the mean  $\mu$  using median polishing. [Cressie, 1993]
- 2. With the deviations  $\delta$  build a parametric covariance function. [Cressie, 1993]
- 3. Compute the conditional mean and covariance at each grid point using the observations and the parametric covariance function.

Kriging is a maximum likelihood estimator of a random field Z modeled by

$$Z(s) = \mu(s) + \delta(s),$$

assuming Gaussian, stationary noises  $\delta(s) \sim \mathcal{N}(0, C(s, s))$ .

The steps of kriging:

- 1. Estimate the mean  $\mu$  using median polishing. [Cressie, 1993]
- 2. With the deviations  $\delta$  build a parametric covariance function. [Cressie, 1993]
- 3. Compute the conditional mean and covariance at each grid point using the observations and the parametric covariance function.

We compare ordinary kriging with a deterministic linear interpolation.

# Spatial Interpolation Results

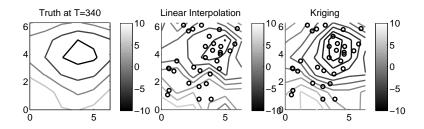


Figure: The true field (left) and the results of a linear interpolation (middle) and kriging interpolation (right).

# Spatial Interpolation Results

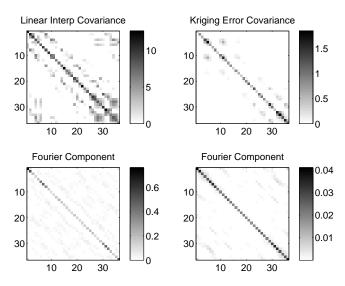


Figure: The noise covariance in physical space (top) and Fourier space (bottom).

The next step is to apply a reduced stochastic Fourier based filter. The filter approximates the barotropic modes of the 2 layer QG model

$$\frac{\partial q}{\partial t} + J(\Psi, q) + \beta \frac{\partial \Psi}{\partial x} + \kappa \nabla^8 q$$

$$+ \left[ J(\Psi^c, q^c) + U \frac{\partial \nabla^2 \Psi^c}{\partial x} - \kappa \nabla^2 \Psi^c \right] = 0$$

in Fourier space with

$$d\hat{\Psi}(t) = (-d + i\omega)\hat{\Psi}(t)dt + Fdt + \sigma dW(t).$$

[Madja and Harlim, Chapter 12, 2012]

The next step is to apply a reduced stochastic Fourier based filter. The filter approximates the barotropic modes of the 2 layer QG model

$$\frac{\partial q}{\partial t} + J(\Psi, q) + \beta \frac{\partial \Psi}{\partial x} + \kappa \nabla^8 q$$

$$+ \left[ J(\Psi^c, q^c) + U \frac{\partial \nabla^2 \Psi^c}{\partial x} - \kappa \nabla^2 \Psi^c \right] = 0$$

in Fourier space with

$$d\hat{\Psi}(t) = (-d + i\omega)\hat{\Psi}(t)dt + Fdt + \sigma dW(t).$$

 $\hat{\Psi}$ : the horizontal Fourier component of the barotropic streamfunction  $\Psi$ ,

[Madja and Harlim, Chapter 12, 2012]



The next step is to apply a reduced stochastic Fourier based filter. The filter approximates the barotropic modes of the 2 layer QG model

$$\frac{\partial q}{\partial t} + J(\Psi, q) + \beta \frac{\partial \Psi}{\partial x} + \kappa \nabla^8 q$$

$$+ \left[ J(\Psi^c, q^c) + U \frac{\partial \nabla^2 \Psi^c}{\partial x} - \kappa \nabla^2 \Psi^c \right] = 0$$

in Fourier space with

$$d\hat{\Psi}(t) = (-d + i\omega)\hat{\Psi}(t)dt + Fdt + \sigma dW(t).$$

 $\hat{\Psi}$ : the horizontal Fourier component of the barotropic streamfunction  $\Psi$ , W(t) : a complex-valued Wiener process,

[Madja and Harlim, Chapter 12, 2012]



The next step is to apply a reduced stochastic Fourier based filter. The filter approximates the barotropic modes of the 2 layer QG model

$$\begin{aligned} \frac{\partial q}{\partial t} + J(\Psi, q) &+ \beta \frac{\partial \Psi}{\partial x} + \kappa \nabla^8 q \\ &+ \left[ J(\Psi^c, q^c) + U \frac{\partial \nabla^2 \Psi^c}{\partial x} - \kappa \nabla^2 \Psi^c \right] = 0 \end{aligned}$$

in Fourier space with

$$d\hat{\Psi}(t) = (-d + i\omega)\hat{\Psi}(t)dt + Fdt + \sigma dW(t).$$

 $\hat{\Psi}$ : the horizontal Fourier component of the barotropic streamfunction  $\Psi$ ,

W(t): a complex-valued Wiener process,

d: damping,

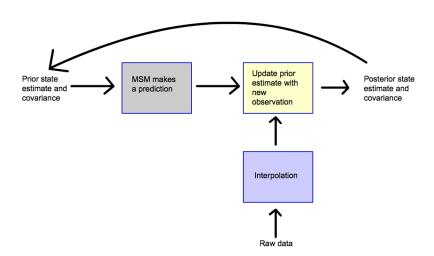
 $\omega$ : frequency,

F: constant external forcing,

 $\sigma$ : noise strength. [Madja and Harlim, Chapter 12, 2012]

#### The Kalman Filter

The Kalman Filter is a solution to these equations and produces estimates of the mean and covariance prior and posterior to observation



#### Filtering Results

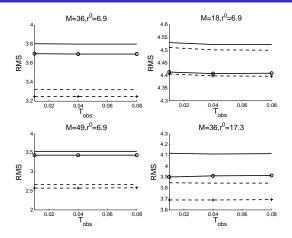


Figure: The RMS errors associated with each step: unfiltered kriging (dashes), filtered kriging (dashes with '+' sign), unfiltered linear interpolation (solid line), and filtered linear interpolation (solid line with circles).

# Filtering Results: M = 36 and $r^0 = 17.3$

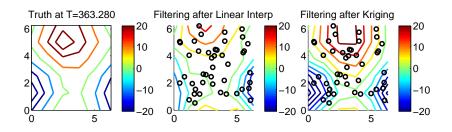


Figure: Filtering results at one particular time. The circles illustrate observation locations.

# Filtering Results: M = 36 and $r^0 = 17.3$

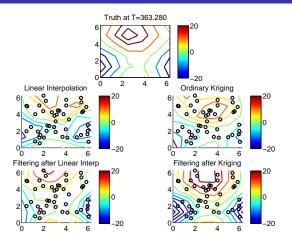


Figure: The true barotropic streamfunction (top), interpolated results (middle panels) and filtered results (bottom panels) at one particular time. The circles illustrate observation locations.

• In every case, kriging outperformed the linear interpolation.

- In every case, kriging outperformed the linear interpolation.
- Filtering further improved the results.

- In every case, kriging outperformed the linear interpolation.
- Filtering further improved the results.
- However the biggest improvements occurred in the cases of sparser observations or larger noise.

- In every case, kriging outperformed the linear interpolation.
- Filtering further improved the results.
- However the biggest improvements occurred in the cases of sparser observations or larger noise.
- The Mean Stochastic Model is a very simple one, and we expect the results could be improved with other models.

#### References

- K. Brown, J. Harlim, Assimilating irregularly spaced sparsely observed turbulent signals with hierarchical Bayesian reduced stochastic filters, Journal of Computational Physics (In Press).
- M. Chanine, Remote Sounding of Cloudy Atmospheres. I. The Single Cloud Layer, American Meteorological Society, 1973.
- N. Cressie, Statistics for Spatial Data, Wiley-Interscience, rev sub edition, 1993.
- J. Harlim, A. Majda, Filtering turbulent sparsely observed geophysical flows, Monthly Weather Review 138 (2010) 1050-1083.
- A. Majda, J. Harlim, Flltering Complex Turbulent Systems, Cambridge University Press, UK, 2012.